

**มหาวิทยาลัยสงขลานครินทร์**  
**คณะวิศวกรรมศาสตร์**

การสอบกลางภาค ประจำภาคการศึกษาที่ 2

ปีการศึกษา 2554

สอบวันที่ 20 ธันวาคม 2554

เวลา 9:00-12:00 น

วิชา 220-622 Groundwater Flow Modeling

ห้องสอบ Robot

**ข้อกำหนด**

1. ข้อสอบ มี 4 ข้อ จำนวน 5 หน้า คะแนนเต็ม 200 คะแนน ให้ทำทุกข้อ
2. อนุญาตให้ Lecture note เข้าห้องสอบได้
3. ให้นำเครื่องคิดเลขทุกชนิดเข้าห้องสอบได้
4. ให้นักศึกษาตอบในสมุดคำตอบ

ชื่อ.....รหัส.....

ข้อ	คะแนนเต็ม	คะแนนที่ได้
1	40	
2	60	
3	50	
4	50	
Total	200	

ออกข้อสอบโดย อ. ธนิต เจริมยานนท์

14 ธ.ค. 2554

## 1. Definitions

Briefly define and explain terms shown in each question. Each definition is worth 5 points.

- 1.1 Central difference approximation
- 1.2 Laplace's equation
- 1.3 Specific storage
- 1.4 Successive Over-Relaxation
- 1.5 Explicit finite difference approximation
- 1.6 Conceptual model
- 1.7 Dupuit's assumptions
- 1.8 Dirichlet and Neumann conditions

## 2. Equations and Finite Difference Approximation

- 2.1 (10 points) Write the governing equation for two-dimensional, steady state flow in homogeneous, but anisotropic unconfined aquifer under Dupuit assumptions. Include a source term to simulate recharge to the aquifer.
- 2.2 (15 points) The mass balance equation for flow through an representative elementary volume is

$$\text{div } q = 0$$

- (a) Write this equation in three dimensions using partial differentials.
  - (b) Write Darcy's law in vector notation and combine it with  $\text{div } q = 0$  to derive the general governing equation groundwater flow. What is the name of the equation just derived.
- 2.3 (20 points) Write a finite difference expression for the following equation, where  $D_x$ ,  $D_y$ ,  $v$  are constant but  $\Delta x \neq \Delta y$

$$D_x \frac{\partial^2 c}{\partial x^2} + D_y \frac{\partial^2 c}{\partial y^2} - v \frac{\partial c}{\partial x} = 0$$

Also write finite difference formula that  $c_{i,j}$  can be computed for the specific case where  $D_x = D_y = D$  and  $\Delta x = \Delta y = a$

2.4 (15 points) From conservation of mass, derive governing equation of transient flow for 2-D unconfined aquifer with sink/source term and also write explicit finite difference approximation for this equation.

3. System Conceptualization

3.1 (25 points) Develop a mathematical model (governing equation and boundary conditions) for the flow system illustrated in the sketch below. Indicate the locations of the boundary conditions mathematically rather than by labeling a sketch. Note that R and q must be treated as boundary conditions.

A sketch shows a profile view of a heterogeneous, anisotropic unconfined aquifer under steady state conditions. The q is outflow flux.

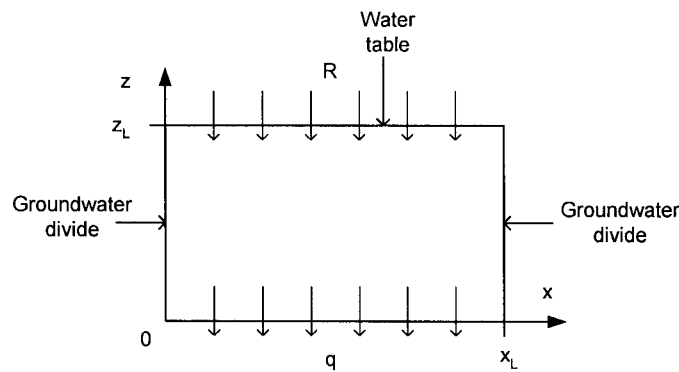


Fig.1 An aquifer system for problem 3.1

3.2 (25 points) The unconfined island recharge problem, with addition of a newly installed pumping well at location B as illustrated in the sketch below. Sea water level is 100 m from datum.

(a) Write the governing equation that you would use to solve the transient problem to predict the decline in head as a result of pumping.

- (b) How would you generate the initial conditions in term of head for transient problem.
- (c) Write the equation for computing a water balance in terms of inflow, outflow and storage for this transient problem.

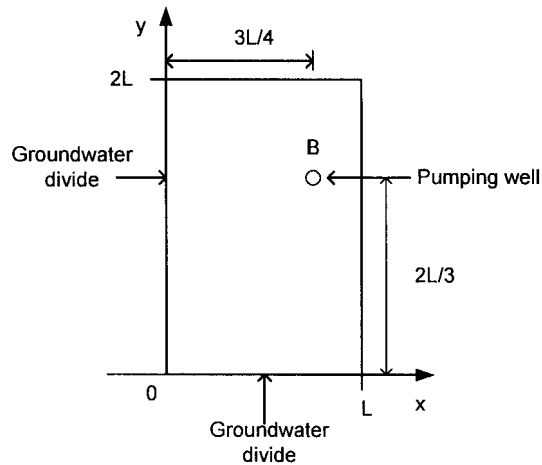


Fig. 2 Island recharge problem for problem 3.2

#### 4. Head and Flux Determination

- 4.1 (25 points) Consider the 6 nodes shown below. Suppose these nodes are the 6 nodes in the upper left hand corner of a two-dimensional horizontal finite difference grid of a confined aquifer. The nodes in the top row are specified head cells and the nodes in the bottom row are active cells in a steady-state simulation. The numbers in each cell are the head values (block-centered) . Suppose we wish to switch from specified head to specified flux boundary conditions. Calculate the values of the flux ( $\text{m}^3/\text{day}$ ) that you would assign to boundary cell A and B. Be sure to indicate the sign of the flux, where (+) indicates inflow to the model and (-) indicates withdrawal of water from the model. Assume that the aquifer is isotropic and transmissivity is  $300 \text{ m}^2/\text{day}$ , and the nodes are equally spaced such that  $\Delta x = \Delta y = 100 \text{ m}$ .

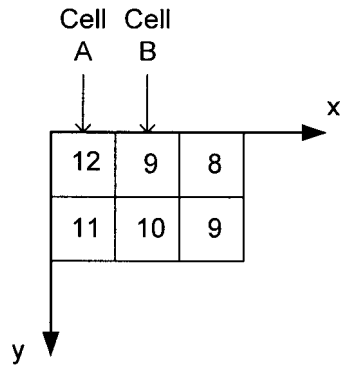


Fig.3 Domain for flux calculation (numbers in the cell indicating the head at each cell)

- 4.2 (25 points) A groundwater flow domain where observed heads (in meters) are shown for each node. Suppose this domain is a 2-D steady-state confined aquifer where Laplace's equation is applied. Calculate the heads at node A and B.

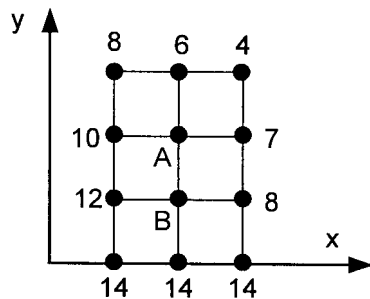


Fig. 4 A confined aquifer domain